

Derivation of Delay Equation Climate Models using Projection Methods

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Introduction

- Conceptual climate models
 - Study physical mechanisms of climate variability
- Differential delay models
 - Infinite-dimensional, but can be formulated in terms of a single variable
 - Mostly introduced in an ad-hoc manner
- Projection methods can place derivation on stronger mathematical foundation
 - Mori-Zwanzig Formalism

Introduction

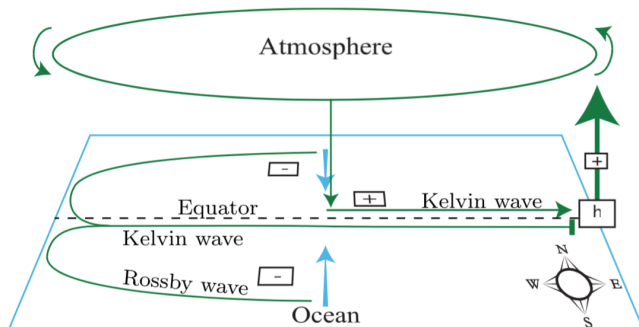
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- Projection methods can place derivation on stronger mathematical foundation
 - Mori-Zwanzig Formalism
- El Niño Southern Oscillation (ENSO)

Delay Model of ENSO (Suarez and Schopf (1988))

$$\frac{dT_e}{dt} = T_e(t) - T_e^3(t) - \alpha T_e(t - \delta)$$

El Niño Southern Oscillation (ENSO)

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¹Keane, Krauskopf, Postlethwaite, 2017.

Mori-Zwanzig Formalism

Linear example for $\phi = (\hat{\phi}, \tilde{\phi}) : \mathbb{R} \rightarrow \mathbb{R}^n$ continuously differentiable:

$$\frac{d}{dt} \begin{pmatrix} \hat{\phi} \\ \tilde{\phi} \end{pmatrix} = \begin{pmatrix} A_{11} & A_{12} \\ A_{21} & A_{22} \end{pmatrix} \begin{pmatrix} \hat{\phi} \\ \tilde{\phi} \end{pmatrix}, \quad \begin{pmatrix} \hat{\phi}(0) \\ \tilde{\phi}(0) \end{pmatrix} = \begin{pmatrix} \hat{x} \\ \tilde{x} \end{pmatrix}.$$

Goal: Equation for *resolved* variables $\hat{\phi} \in \mathbb{R}^m$ only,
the *unresolved* variables are $\tilde{\phi} \in \mathbb{R}^{n-m}$.

Mori-Zwanzig Formalism

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$$\frac{d}{dt} \hat{\phi}(t) = A_{11} \hat{\phi}(t) + A_{12} e^{A_{22}t} \tilde{x} + \int_0^t A_{12} e^{A_{22}(t-s)} A_{21} \hat{\phi}(s) ds$$

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Markovian

Noise

Memory

Mori-Zwanzig Formalism

In general, consider a **system of ODEs**:

$$\frac{d}{dt}\phi(t) = R(\phi(t)), \quad \phi(0) = x,$$

$\phi(t) \in \mathbb{R}^n$ continuously differentiable, $R : \mathbb{R}^n \rightarrow \mathbb{R}^n$.

Evolution of an observable $u(x, t) := g(\phi(x, t))$ along a solution satisfies the PDE

$$\frac{\partial}{\partial t}u(x, t) = \mathcal{L}u(x, t), \quad u(x, 0) = g(x),$$

where $[\mathcal{L}u](x) = \sum_{i=1}^n R_i(x)\partial_{x_i}u(x)$ is the **Liouville operator**.

A **projection** P onto a set of resolved variables $\hat{\phi}$, with complement $Q = I - P$.

Mori-Zwanzig Formalism

$$\frac{d}{dt} \hat{\phi}(t) = A_{11} \hat{\phi}(t) + A_{12} e^{A_{22}t} \tilde{X} + \int_0^t A_{12} e^{A_{22}(t-s)} A_{21} \hat{\phi}(s) ds$$

Generalized Langevin Equation²

$$\frac{\partial}{\partial t} \phi_i(x, t) = R_i(\hat{\phi}(x, t)) + F_i(x, t) + \int_0^t K_i(\hat{\phi}(x, t-s), s) ds,$$

with

$$F_i(x, t) = [e^{tQ\mathcal{L}} Q\mathcal{L}g](x), \quad K_i(\hat{x}, t) = [P\mathcal{L}F_i](x, t).$$

$F_i(x, t)$ solves the **orthogonal dynamics equation**:

$$\frac{d}{dt} F_i(x, t) = Q\mathcal{L}F_i(x, t), \quad F_i(x, 0) = Q\mathcal{L}x_i.$$

²For derivation see Chorin, Hald, Kupferman, 2002.

ENSO Model

Two-Strip Model (rewritten)³

$$\partial_t h_c + \epsilon_0 h_c + \partial_x h_c = \mu \left(1 - \frac{\theta}{1 + y_n^2}\right) g(x) T_e(x_E, t)$$

$$\partial_t h_n + \epsilon_0 h_n - \frac{1}{y_n^2} \partial_x h_n = -\mu \frac{\theta}{y_n^2} g(x) T_e(x_E, t)$$

$$\partial_t T_e + c_T T_e - c_h \left(h_c + \frac{1}{1 + y_n^2} h_n \right) = 0$$

T_e	Temperature at equator
h_e	Thermocline at equator
h_n	Thermocline at $y = y_n$

where $h_c(x, t) = h_e(x, t) - \frac{1}{1 + y_n^2} h_n(x, t)$. The boundary conditions are:

$$h_c(0, t) = \left(r_W - \frac{1}{1 + y_n^2} \right) h_n(0, t), \quad h_c(1, t) = \left(\frac{1}{r_E} - \frac{1}{1 + y_n^2} \right) h_n(1, t).$$

³Rewritten from Jin, 1996.

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$$\begin{aligned}
 \partial_t h_c + \epsilon_0 h_c + \boxed{\partial_x h_c} &= \mu \left(1 - \frac{\theta}{1 + y_n^2} \right) g(x) T_e(x_E, t) \\
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 \partial_t T_e + \boxed{c_T} T_e - \boxed{c_h} \left(h_c + \frac{1}{1 + y_n^2} h_n \right) &= 0
 \end{aligned}
 \left| \begin{array}{l} T_e \text{ Temperature} \\ \text{at equator} \\ h_e \text{ Thermocline} \\ \text{at equator} \\ h_n \text{ Thermocline} \\ \text{at } y = y_n \end{array} \right.$$

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Consider a **linear** version, i.e. no dependence of c_T , c_h on e.g. T_e .
Use a linear **projection onto** T_e .

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Mori-Zwanzig Formalism

$$\begin{aligned} \frac{dT_e}{dt}(x, t) = & -c_T(x)T_e(x, t) \\ & + c_h(x) \left(e^{-(\epsilon_0 + \partial_x)t} h_c(x, 0) + \frac{1}{1 + y_n^2} e^{-(\epsilon_0 - \frac{1}{y_n^2} \partial_x)t} h_n(x, 0) \right) \\ & + \int_0^t c_h(x) \left(B_0 e^{-(\epsilon_0 + \partial_x)(t-s)} - B_1 e^{-(\epsilon_0 - \frac{1}{y_n^2} \partial_x)(t-s)} \right) \\ & \quad \cdot g(x) T_e(x_E, s) ds \end{aligned}$$

See derivation in arXiv:1902.03198.

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Is this a delay equation?

Characteristics

Memory-Term

$$\int_0^t c_h(x_E) \left[\left(B_0 e^{-(\epsilon_0 + \partial_x)(t-s)} - B_1 e^{-(\epsilon_0 - \frac{1}{y_n^2} \partial_x)(t-s)} \right) \cdot g(x) \right]_{x_E} T_e(x_E, s) ds$$

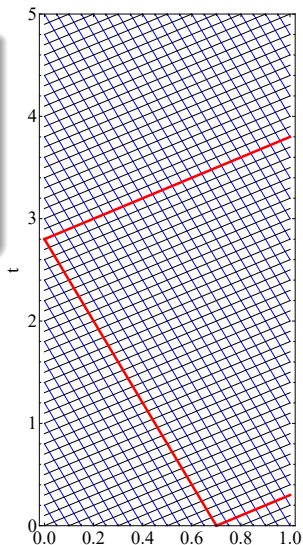
- Interested in east of the basin: $x = x_E$

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- Interested in east of the basin: $x = x_E$
- Follow signal along characteristics to eastern boundary
 - $\partial_t f = -\epsilon_0 f - \partial_x f \rightarrow x - x_0 = t - t_0$
 - $\partial_t f = -\epsilon_0 f + \frac{1}{y_n^2} \partial_x f \rightarrow x - x_0 = \frac{-1}{y_n^2} (t - t_0)$
- No reflection at eastern boundary
 - Energy loss at western boundary: A_{rW}

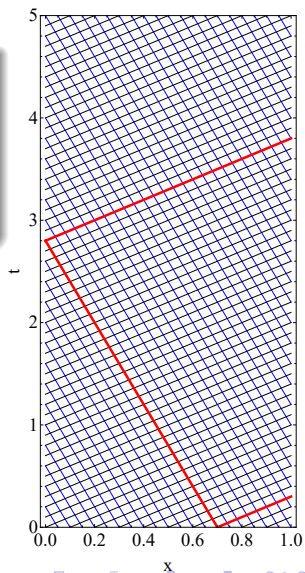


Characteristics

Memory-Term

$$c_h(x_E)A_0 \left(B_0 e^{-\epsilon_0(1-x_w)} T_e(x_E, t - (1 - x_w)) \right. \\ \left. - B_1 A_{rW} e^{-\epsilon_0(1+y_n^2 x_w)} T_e(x_E, t - (1 + y_n^2 x_w)) \right)$$

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- No reflection at eastern boundary
 - Energy loss at western boundary: A_{rW}
- Wind forcing acts locally: $g(x) = A_0 \delta_{x_w}(x)$



Linear Delay Equation

$$\begin{aligned} \frac{dT_e^E}{dt} = & -c_T(x_E)T_e^E(t) \\ & + c_h(x_E)A_0 \left(B_0 e^{-\epsilon_0(1-x_w)} T_e^E(t - (1-x_w)) \right. \\ & \left. - B_1 A_{rW} e^{-\epsilon_0(1+y_n^2 x_w)} T_e^E(t - (1+y_n^2 x_w)) \right) \end{aligned}$$

Note the noise term vanishes by assuming no reflection at the eastern boundary.

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Note the noise term vanishes by assuming no reflection at the eastern boundary.

Since $1 - x_w \ll 1 + y_n^2 x_w$ we assume $T_e^E(t - (1 - x_w)) \approx T_e^E(t)$.

Delay Model ENSO

$$\frac{dT_e^E}{dt} = c_S T_e^E(t) - c_L T_e^E(t - d)$$

Nonlinear ENSO Model

Nonlinear Temperature Equation⁴

$$\partial_t T_e + c_T(x) T_e - c_h^*(x)(1 - \beta T_e^2) \left(h_c + \frac{1}{1 + y_n^2} h_n \right) = 0$$

⁴Based on Dijkstra, Neelin, 1995.

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Two approaches:

- 1 Approximation to Mori-Zwanzig formalism

$$\begin{aligned} \frac{dT_e^E}{dt} = & (c_S^* - c_T(x_E)) T_e^E(t) - c_L^* T_e^E(t-d) - \beta c_S^* T_e^E(t)^3 \\ & + \beta c_L^* T_e^E(t-d)^3 \end{aligned}$$

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- 2 Variation of constants

- Equations for h_c and h_n are still linear

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ENSO Delay Models

Suarez and Schopf Model (S&S)

$$\frac{dT}{dt} = T(t) - T^3(t) - \alpha T(t - \delta)$$

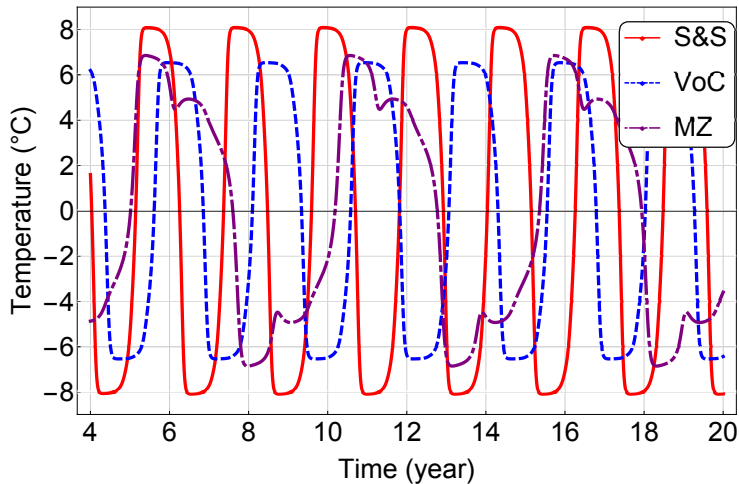
Variation of Constants Model (VoC)

$$\frac{dT}{dt} = T(t) - T^3(t) - \alpha T(t - \delta)(1 - \gamma T^2(t))$$

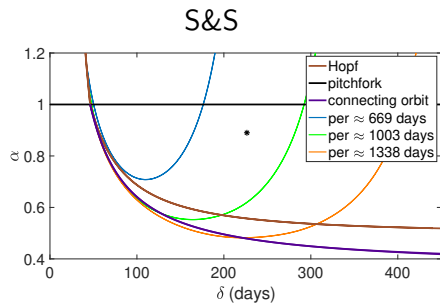
Mori-Zwanzig Model (MZ)

$$\frac{dT}{dt} = T(t) - T^3(t) - \alpha T(t - \delta)(1 - \gamma T^2(t - \delta))$$

Periodic Solutions

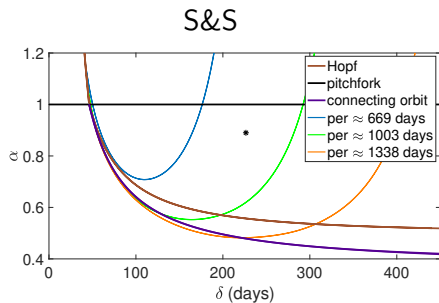


Bifurcation Diagrams

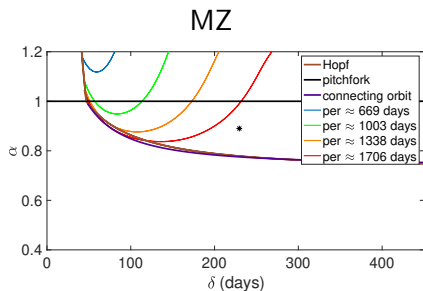
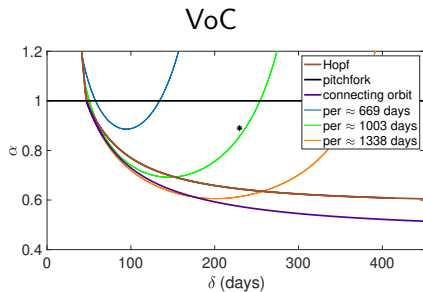


Top: $\gamma = 0$.

Bifurcation Diagrams



Top: $\gamma = 0$. Right: $\gamma = 0.49$.



Summary

- Mori-Zwanzig formalism can be used to derive delay equation models
 - When the equations for the unresolved variables are linear variation of constants is equivalent
- Application to a two-strip ENSO model leads to an improvement in period compared to a previously studied model
- Method can be extended to other wave equations (firstly, to those which are linear in the unresolved variables)
- For nonlinear models better approximation techniques for the orthogonal dynamics are needed

S.K.J. Falkena, C. Quinn, J. Sieber, J. Frank, H.A. Dijkstra, *Derivation of Delay Equation Climate Models Using the Mori-Zwanzig Formalism*, 2019, arXiv:1902.03198, (under review in PRSA).

Thank you!

