

Predictability of Extreme Events

Christian Franzke

Meteorological Institute

Center for Earth System Research and Sustainability

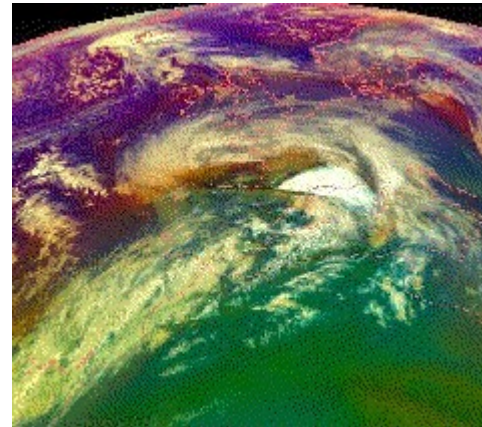
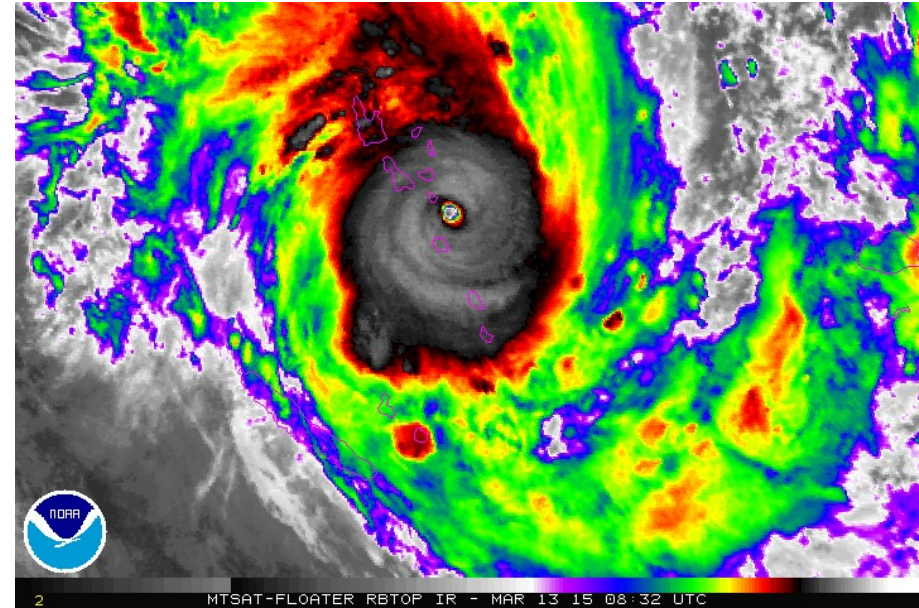
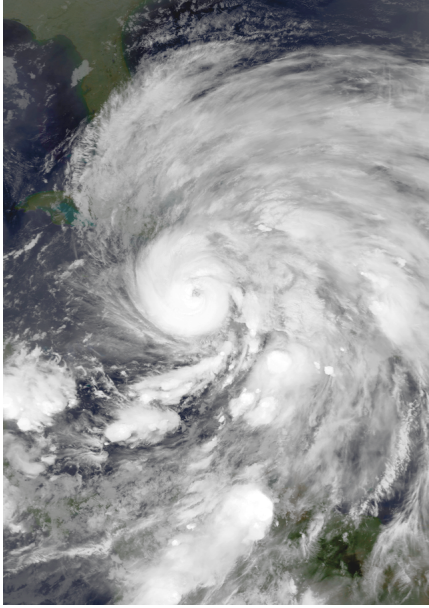
University of Hamburg

In collaboration with **Tamas Bodai** (UHH)

Outline

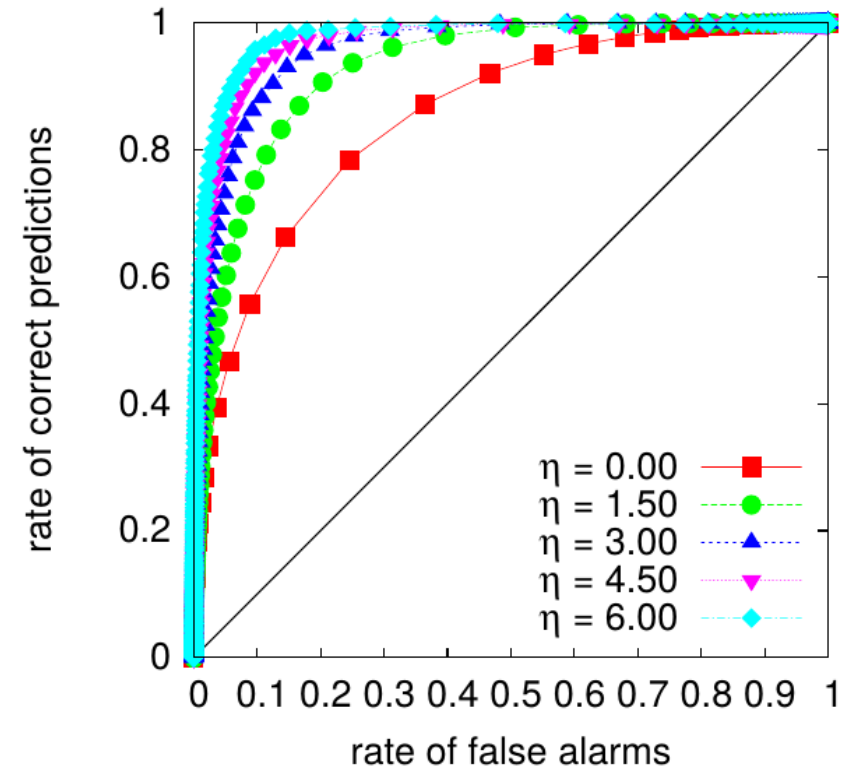
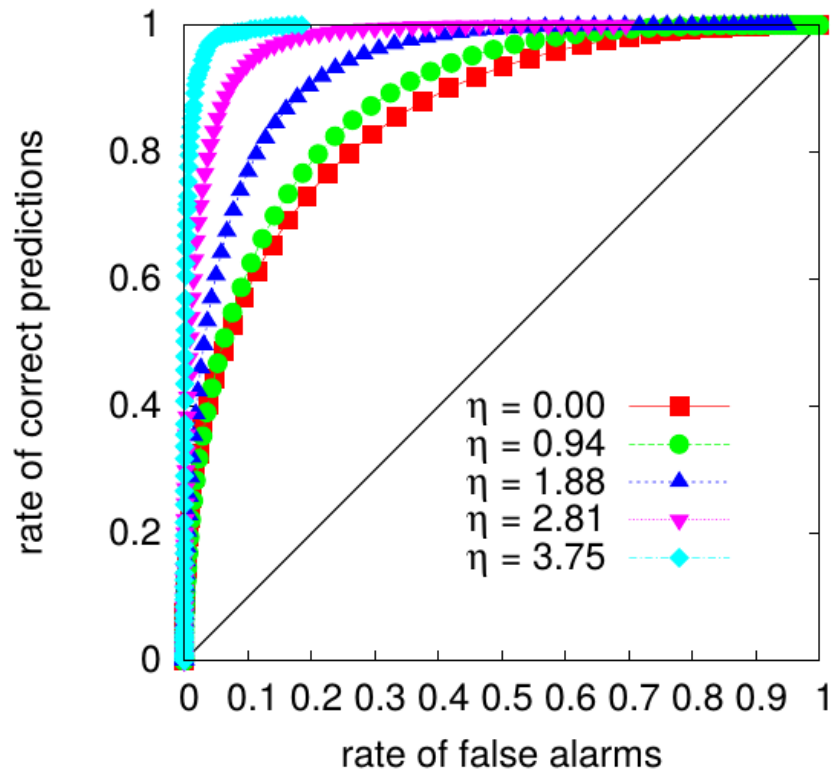
- 1) Motivation
- 2) Dynamic-Stochastic Models for Extremes
- 3) CAM Noise
- 4) Are extremes better predictable the larger they are?
- 5) Summary

Motivation



Motivation

Are extremes better predictable the larger they are?



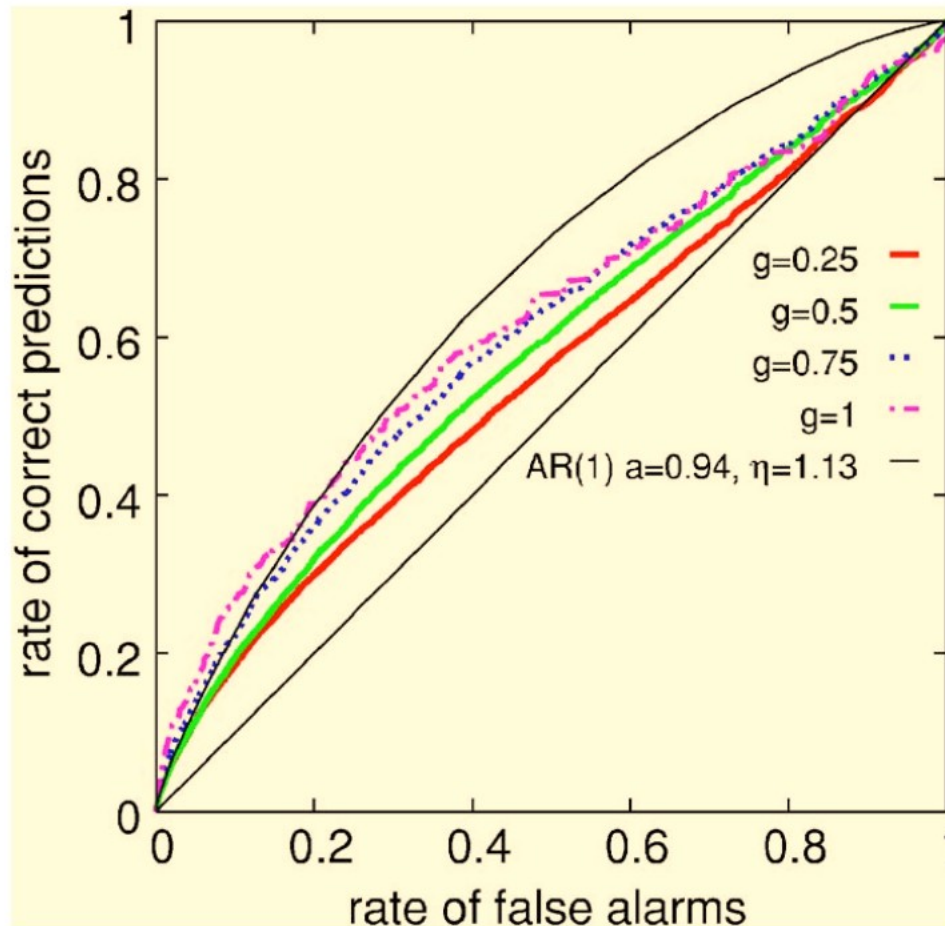
AR(1) with Gaussian and power-law distributed noise

Hallerberg and Kantz 2008

Motivation

Are extremes better predictable the larger they are?

Wind gust data



Hallerberg et al. 2007

How to model extremes?

- Classical Extreme Value Theory (Coles 2001)
- Dynamical Systems Theory of Extremes (Lucarini et al. 2016)
- Stochastic Models of Extremes

Dynamic-Stochastic Modeling

Fundamental form of weather and climate models:

$$dz = \mathbf{F}dt + Lzdt + B(\mathbf{z},\mathbf{z})dt, \quad \mathbf{z} = \begin{pmatrix} \mathbf{x} \\ \mathbf{y} \end{pmatrix},$$

\mathbf{x} : Slow mode
 \mathbf{y} : Fast mode

Dynamic-Stochastic Modeling

Fundamental form of weather and climate models:

$$dz = \mathbf{F}dt + Lzdt + B(\mathbf{z},\mathbf{z})dt, \quad \mathbf{z} = \begin{pmatrix} \mathbf{x} \\ \mathbf{y} \end{pmatrix},$$

Fast modes:

$$d\mathbf{y} = \mathbf{F}_y d\mathbf{t} + L_y \mathbf{y} d\mathbf{t} + L_x \mathbf{x} d\mathbf{t} + B_{yxx}(\mathbf{x}, \mathbf{x}) d\mathbf{t} \\ + B_{yxy}(\mathbf{x}, \mathbf{y}) d\mathbf{t} + B_{yyy}(\mathbf{y}, \mathbf{y}) d\mathbf{t}$$

Dynamic-Stochastic Modeling

Fundamental form of weather and climate models:

$$dz = \mathbf{F}dt + Lzdt + B(\mathbf{z}, \mathbf{z})dt, \quad \mathbf{z} = \begin{pmatrix} \mathbf{x} \\ \mathbf{y} \end{pmatrix},$$

Fast modes:

$$d\mathbf{y} = \mathbf{F}_y d\mathbf{t} + L_y \mathbf{y} d\mathbf{t} + L_x \mathbf{x} d\mathbf{t} + B_{yxx}(\mathbf{x}, \mathbf{x}) d\mathbf{t} \\ + B_{yxy}(\mathbf{x}, \mathbf{y}) d\mathbf{t} + B_{yyy}(\mathbf{y}, \mathbf{y}) d\mathbf{t}$$

Stochastic Modeling Assumption:

$$B_{yyy}(\mathbf{y}, \mathbf{y}) d\mathbf{t} \sim -\frac{\gamma}{\varepsilon} \mathbf{y} + \frac{\sigma}{\sqrt{\varepsilon}} d\mathbf{W}.$$

Dynamic-Stochastic Modeling

Fundamental form of weather and climate models:

$$dz = \mathbf{F}dt + Lzdt + B(\mathbf{z}, \mathbf{z})dt, \quad \mathbf{z} = \begin{pmatrix} \mathbf{x} \\ \mathbf{y} \end{pmatrix},$$

Stochastic Climate Models:

$$d\mathbf{x} = \tilde{\mathbf{F}}dt + \tilde{\mathbf{L}}\mathbf{x}dt + \tilde{\mathbf{B}}(\mathbf{x}, \mathbf{x})dt + \tilde{\mathbf{M}}(\mathbf{x}, \mathbf{x}, \mathbf{x})dt \\ + \tilde{\sigma}_1 dW_1 + \tilde{\sigma}_2(\mathbf{x})dW_2.$$

$$\tilde{\sigma}_2(\mathbf{x})dW_2 = (a + bx_2)dW_2,$$

Dynamic-Stochastic Modeling

Fundamental form of weather and climate models:

$$dz = \mathbf{F}dt + Lzdt + B(\mathbf{z}, \mathbf{z})dt, \quad \mathbf{z} = \begin{pmatrix} \mathbf{x} \\ \mathbf{y} \end{pmatrix},$$

Stochastic Climate Models:

$$d\mathbf{x} = \tilde{\mathbf{F}}dt + \tilde{\mathbf{L}}\mathbf{x}dt + \tilde{\mathbf{B}}(\mathbf{x}, \mathbf{x})dt + \tilde{\mathbf{M}}(\mathbf{x}, \mathbf{x}, \mathbf{x})dt \\ + \tilde{\sigma}_1 dW_1 + \tilde{\sigma}_2(\mathbf{x})dW_2.$$

$$\tilde{\sigma}_2(\mathbf{x})dW_2 = (a + bx_2)dW_2,$$

Correlated Additive Multiplicative Noise
(CAM Noise)

CAM Noise

Normal form of stochastic climate models:

$$\frac{dx}{dt} = F + ax + bx^2 - cx^3 + \sum_p \frac{\sigma_p}{\gamma_p} (L_{1p} - I_{1p}^M x) \dot{W}_p + \sigma_A \dot{W}_A$$

Corresponding Fokker-Planck Equation:

$$\frac{\partial p}{\partial t} = -\frac{\partial}{\partial x} [(F + ax + bx^2 - cx^3)p] + \frac{1}{2} \frac{\partial^2}{\partial x^2} \left[\left(\sum_p \left((L_{1p} - I_{1p}^M x) \frac{\sigma_p}{\gamma_p} \right)^2 + \sigma_A^2 \right) p \right]$$

Majda et al. 2009

CAM Noise

Stationary distribution:

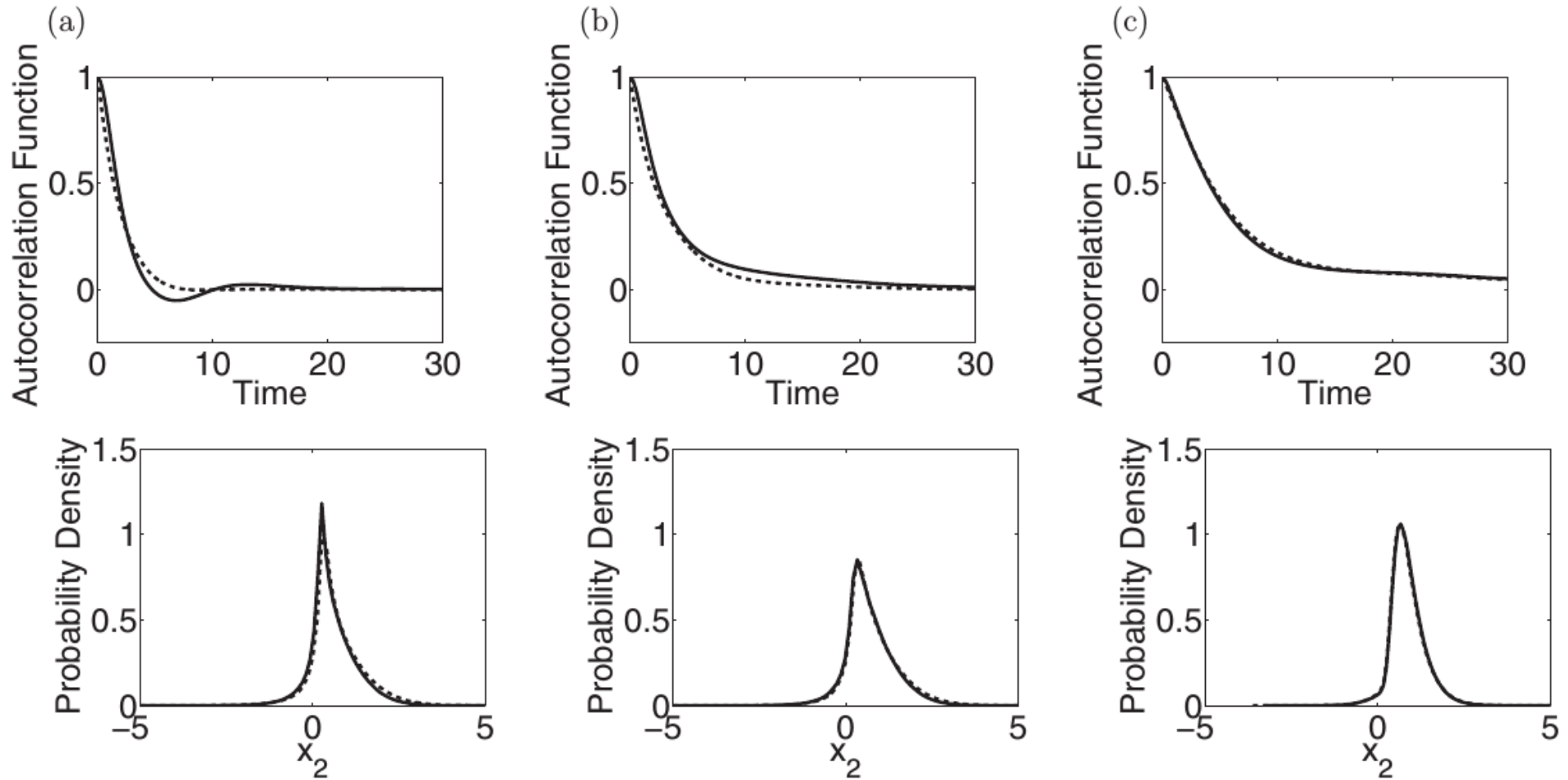
$$p_{\text{Stat}} = \frac{N_0}{((\tilde{A} - \tilde{B}x)^2 + g^2)^{\tilde{a}}} \\ \times \exp\left(\tilde{d} \arctan\left(\frac{\tilde{B}x - \tilde{A}}{g}\right)\right) \\ \times \exp\left(\frac{1}{\tilde{B}^4} (-\tilde{c}x^2 + \tilde{b}x)\right)$$

Power law component

Exponential component

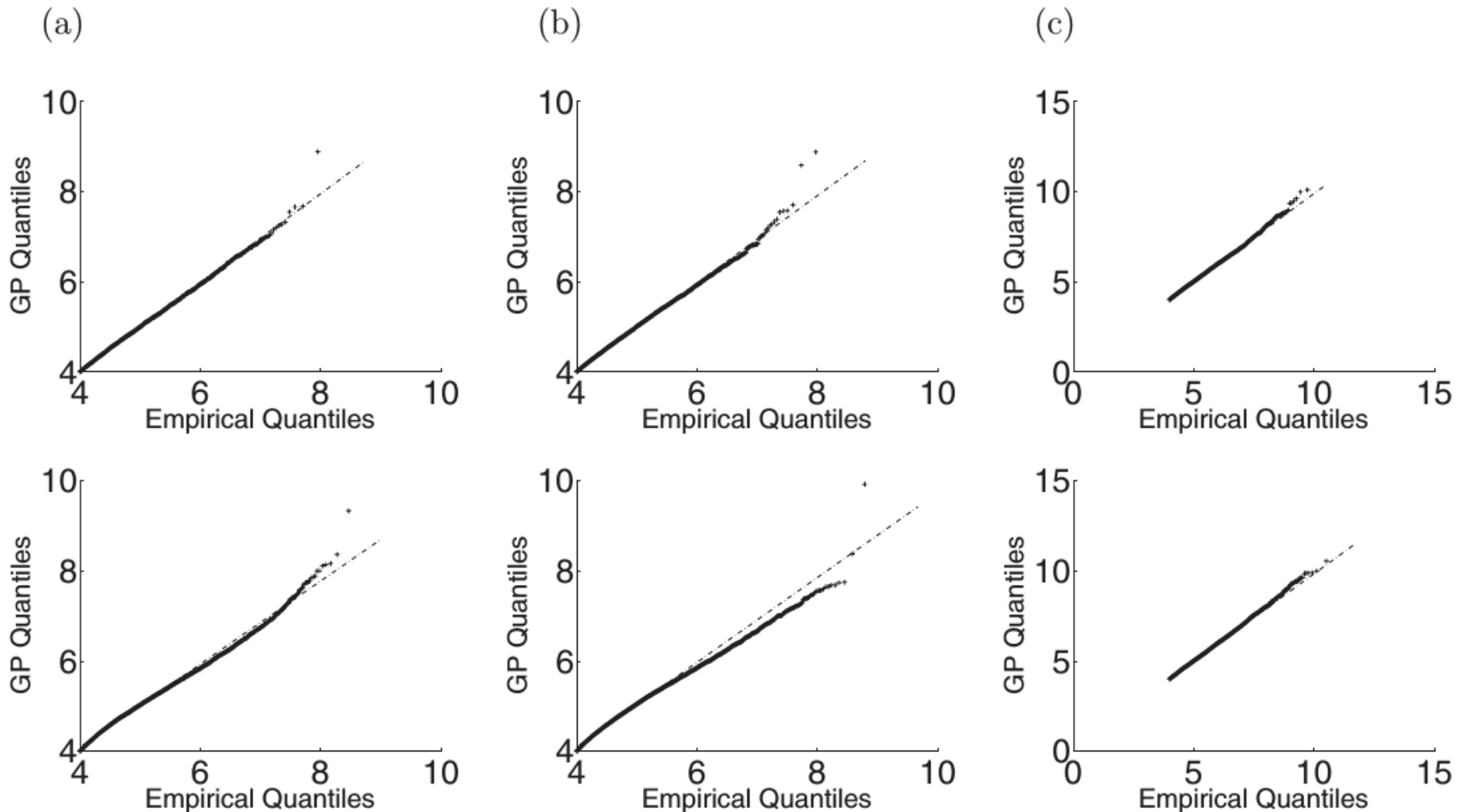
Majda et al. 2009

Conceptual Climate Model



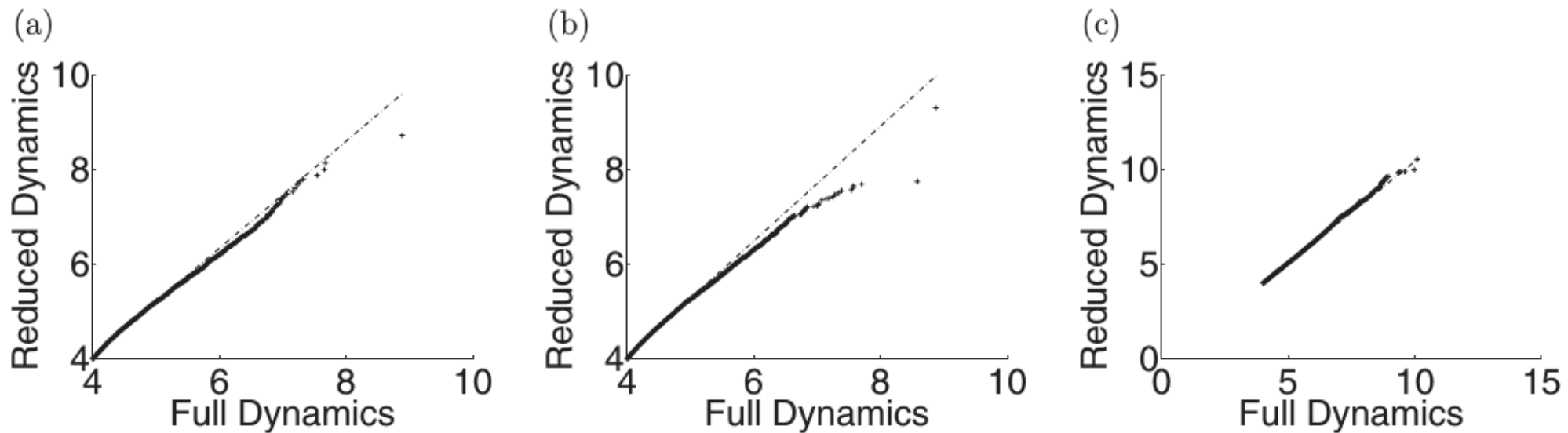
(a) $\varepsilon = 1.0$, (b) $\varepsilon = 0.5$, and (c) $\varepsilon = 0.1$.

Conceptual Climate Model



. (a) $\varepsilon = 1.0$, (b) $\varepsilon = 0.5$, and (c) $\varepsilon = 0.1$.

Conceptual Climate Model



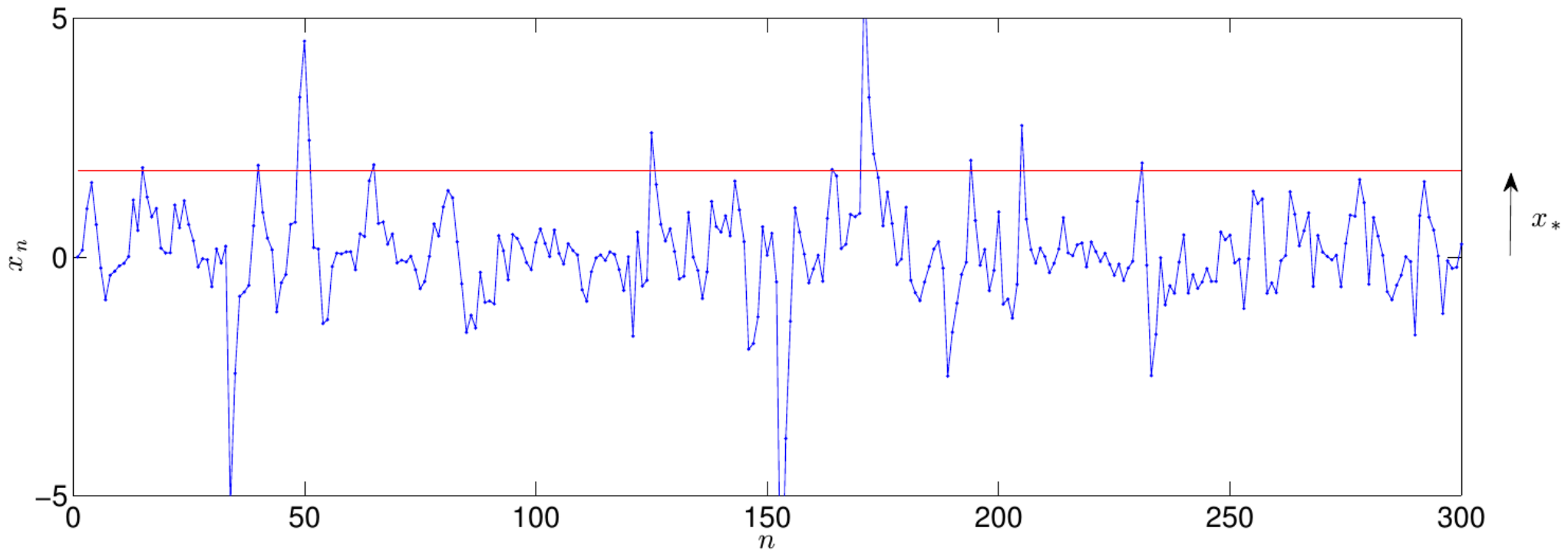
. (a) $\epsilon = 1.0$, (b) $\epsilon = 0.5$, and (c) $\epsilon = 0.1$.

Conceptual Climate Model

Are extremes better predictable the larger they are?

(Hallerberg et al. 2007, 2008; Franzke 2012;
Miotto and Altmann 2014)

Predictability



Prediction Scheme

Binary event variable:

$$\chi_n = \begin{cases} 1, & x(t_n) > x_* \\ 0, & x(t_n) < x_* \end{cases}$$

Binary prediction at $t(n-d)$ is defined as:

$$\hat{\chi}_n = \begin{cases} 1, & \mathcal{L}(\mathbf{x}_n) > \mathcal{L}_* \\ 0, & \mathcal{L}(\mathbf{x}_n) < \mathcal{L}_* \end{cases}$$

Hallerberg et al. 2008, Bodai 2015

Prediction Scheme

Hit rate

$$H(\mathcal{L}_*) = \frac{\int_{\mathbb{R}^M} dV_{\mathbf{x}} \mathcal{P}(\mathbf{x}) \mathcal{H}(\mathcal{L}(\mathbf{x}) - \mathcal{L}_*)}{\int_{\mathbb{R}^M} dV_{\mathbf{x}} \mathcal{P}(\mathbf{x})}.$$

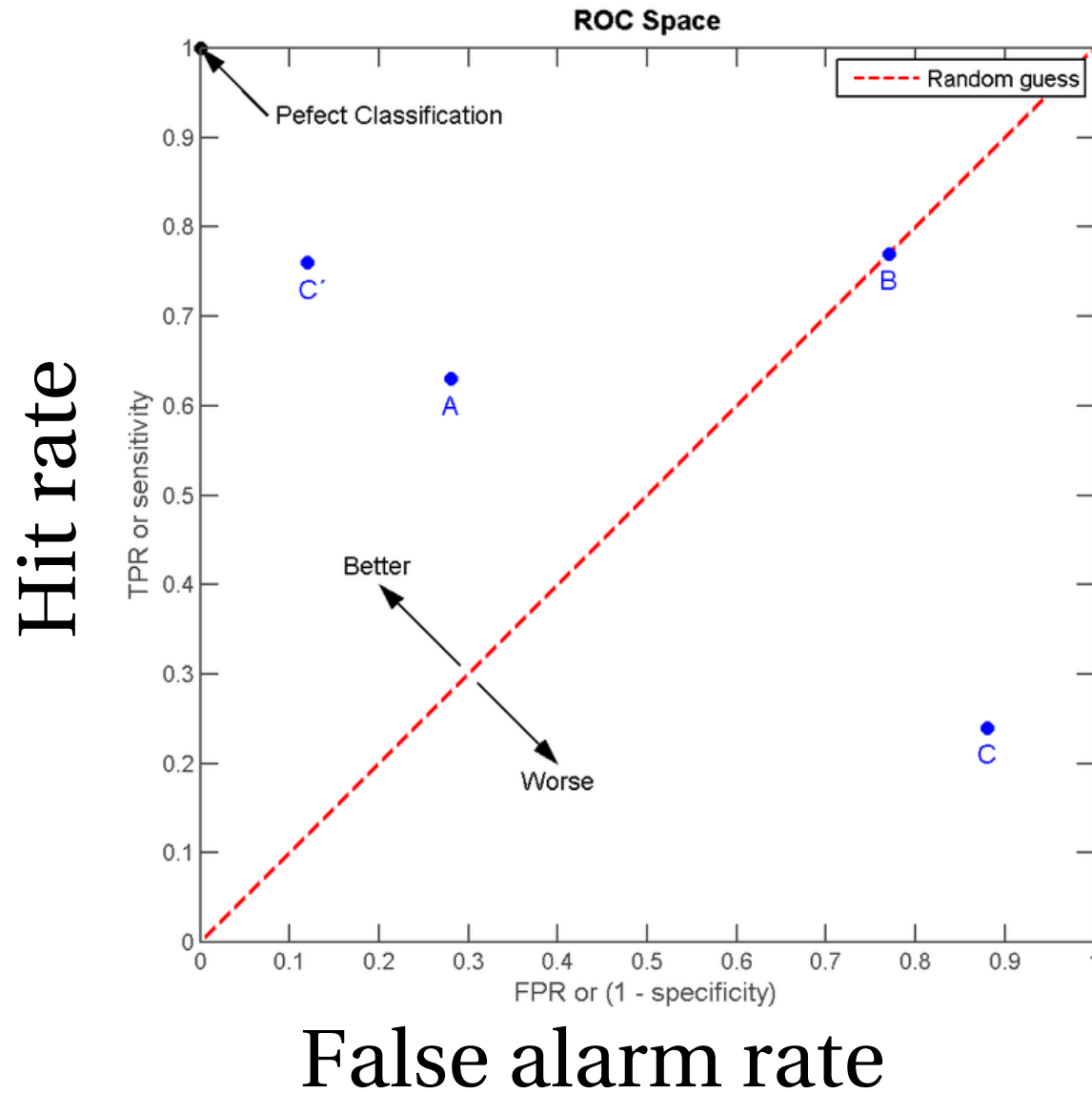
False alarm rate

$$F(\mathcal{L}_*) = \frac{\int_{\mathbb{R}^M} dV_{\mathbf{x}} [p(\mathbf{x}) - \mathcal{P}(\mathbf{x})] \mathcal{H}(\mathcal{L}(\mathbf{x}) - \mathcal{L}_*)}{\int_{\mathbb{R}^M} dV_{\mathbf{x}} [p(\mathbf{x}) - \mathcal{P}(\mathbf{x})]}.$$

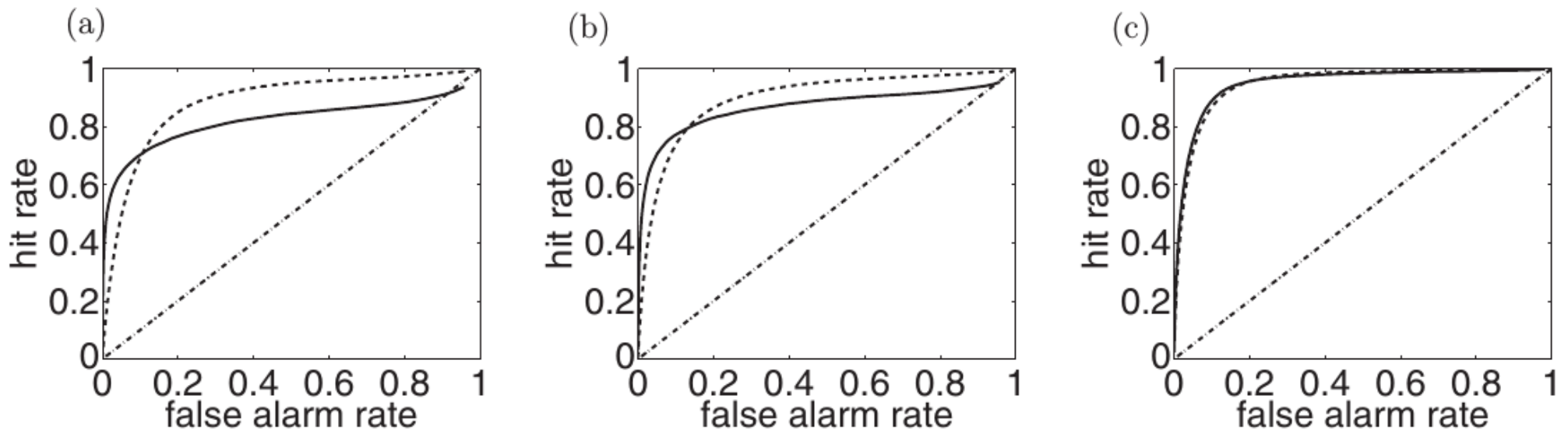
$\mathcal{H}(\cdot)$ is the Heaviside step function

Hallerberg et al. 2008, Bodai 2015

ROC Curves

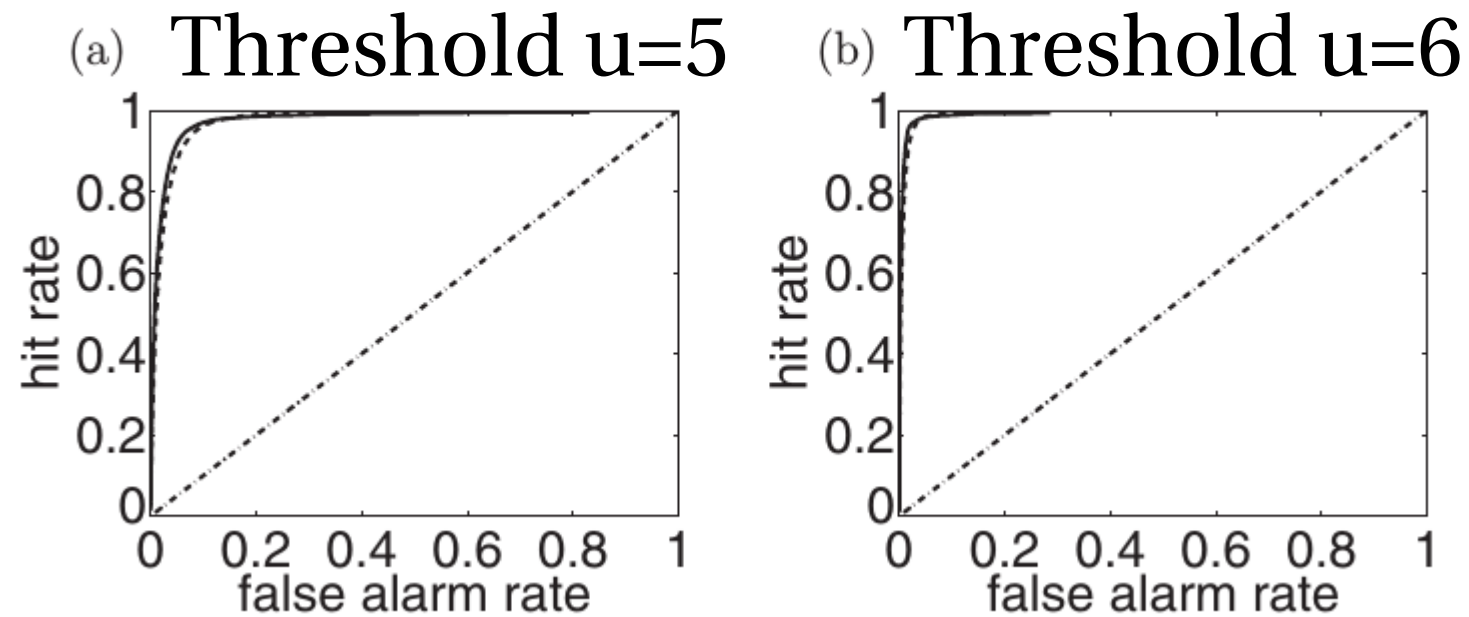


Predictability of Extremes



. (a) $\varepsilon = 1.0$, (b) $\varepsilon = 0.5$, and (c) $\varepsilon = 0.1$.

Predictability of Extremes



Predictability of Extremes

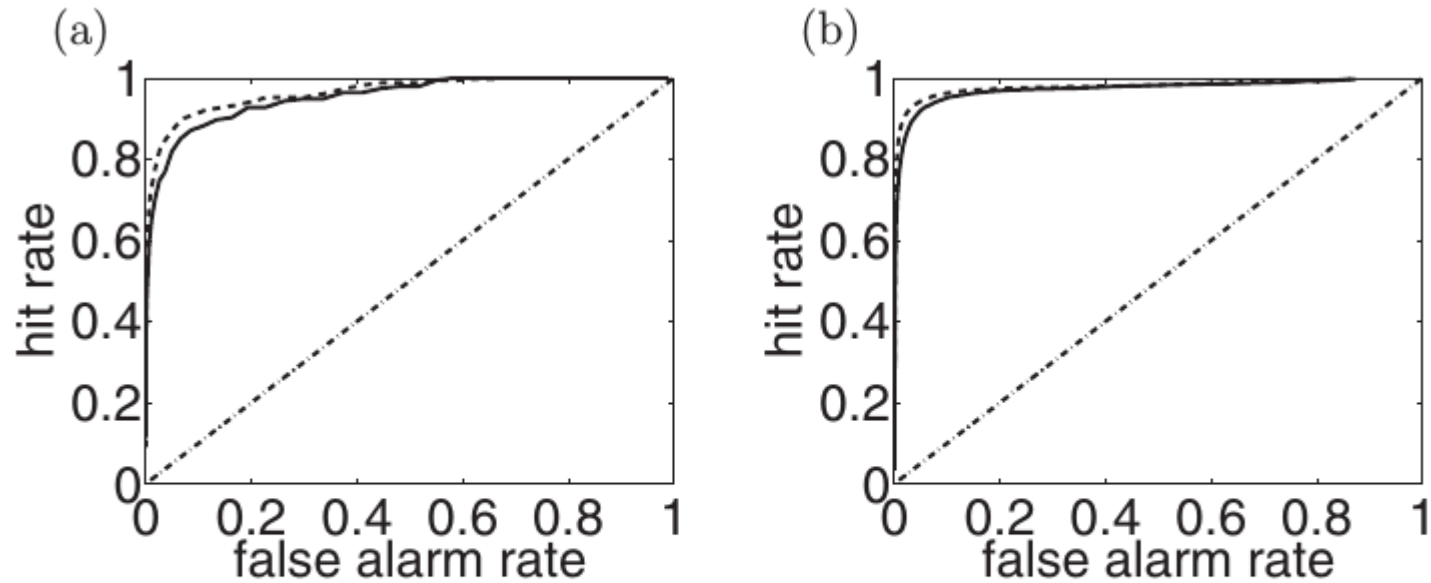


FIG. 9. Receiver-operator characteristic curves: (a) model simulation with no multiplicative triads active and (b) model simulation with no bare truncation active (case with $\varepsilon = 0.1$).

Process Models

AR(1) model with Pareto noise

$$X_{n+1} = aX_n + \xi_n,$$

$$p_\xi(\xi) = p_{sP}(\xi; \alpha, \xi_m) = \alpha \xi_m^\alpha / |\xi + \xi_m|^{\alpha+1}$$

$$\alpha \geq 2 \longrightarrow p(x) \approx N(0, \sigma) = \frac{1}{\sqrt{2\pi}\sigma} e^{-\frac{x^2}{2\sigma^2}}, \quad \sigma = \sigma_\xi / \sqrt{1 - a^2}.$$

$$\alpha < 2 \longrightarrow \sigma_\xi \text{ does not exist,}$$

Process Models

AR(1) model with Pareto noise

$$X_{n+1} = aX_n + \xi_n,$$

$$p_\xi(\xi) = p_{SP}(\xi; \alpha, \xi_m) = \alpha \xi_m^\alpha / |\xi + \xi_m|^{\alpha+1}$$

Process Models

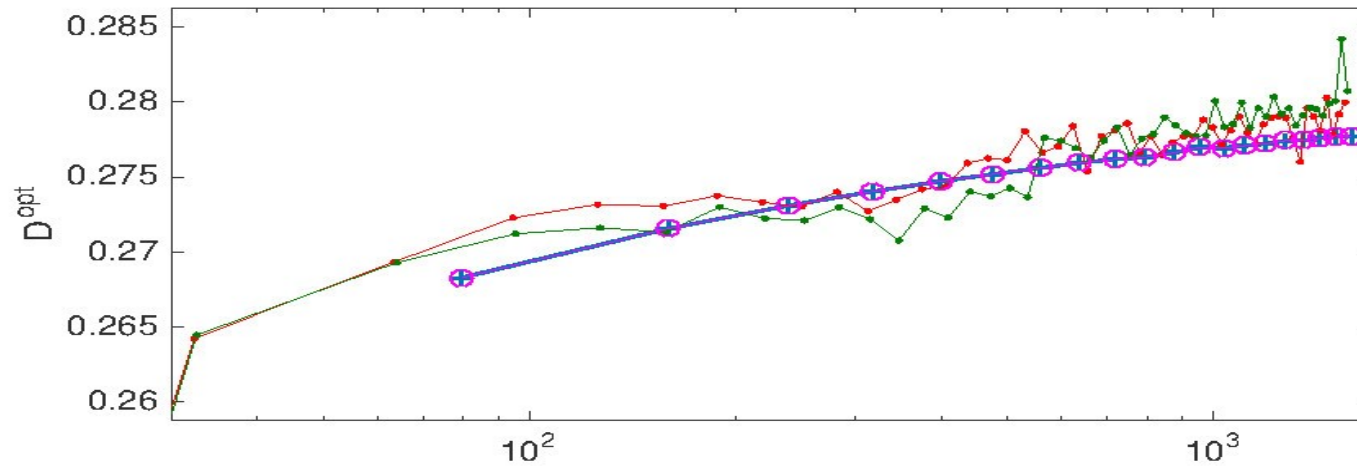
CAM model with α -stable noise

$$dx = (b + ax)dt + (d + cx)dW$$

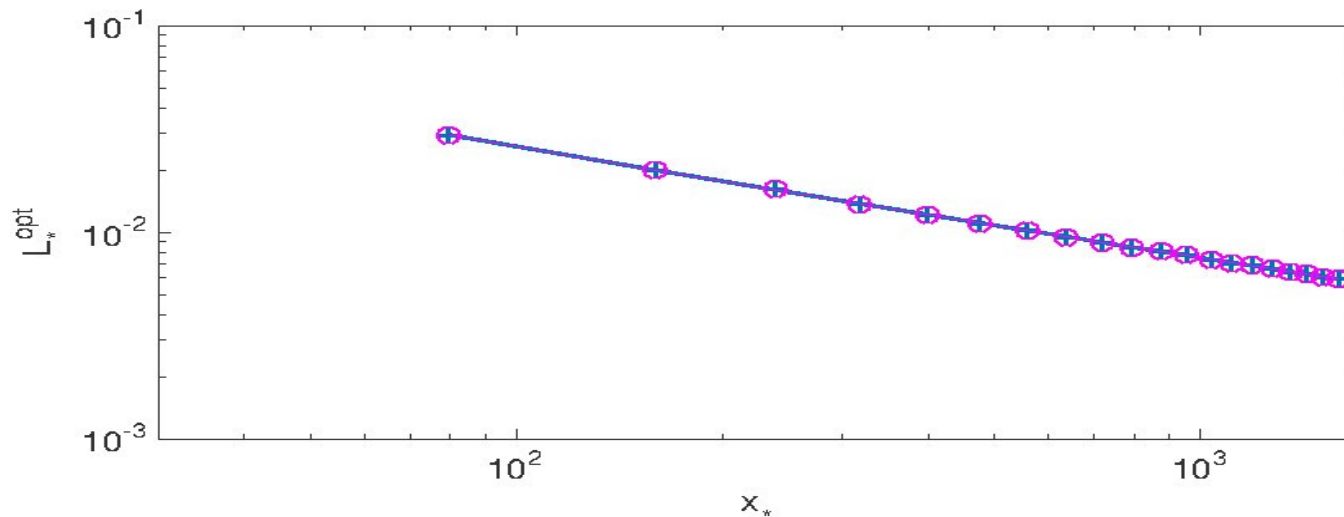
Solution of the stationary Fokker-Planck equation:

$$p(x) = N_0 \frac{2e^{2\frac{ad-bc}{c^2(d+cx)}}}{(d+cx)^{2(1-a/c^2)}}$$

Process Models



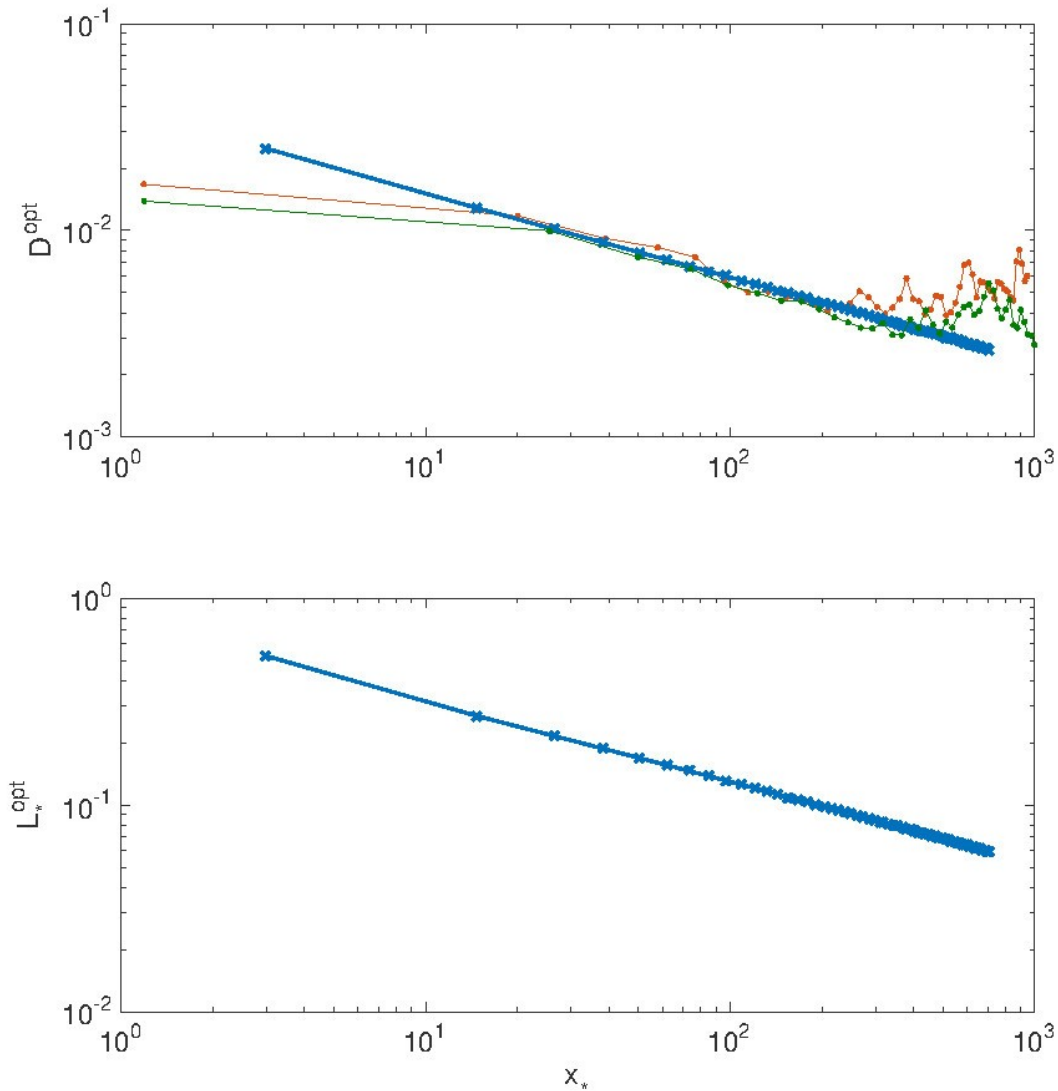
AR(1) model
with
Pareto noise



Distance (for categorical prediction):
$$D = \min_{\mathcal{L}_*} (\sqrt{F^2 + (H - 1)^2})$$

Process Models

CAM model
with α -stable noise



Summary

- Dynamic-Stochastic models have predictive skill when forecasting extreme events
- Extreme events are not necessarily better predictable the more extreme they are
- Predictability seems to be determined by the process

References:

- Majda, Franzke and Crommelin, 2009: Normal forms for reduced stochastic climate models. Proc. Natl. Acad. Sci. USA, 106, 3649-3653.
- Franzke, C., 2012: Predictability of Extreme Events in a Nonlinear Stochastic-Dynamical Model. Phys. Rev. E, 85, DOI: 10.1103/PhysRevE.85.031134.
- Franzke, C., 2016: Extremes in dynamic-stochastic models, Chaos.
- Bodai, T. and C. Franzke, 2016: Predictability of extremes, in preparation.